

Available online at [www.sciencedirect.com](http://www.sciencedirect.com)

ScienceDirect

[www.elsevier.com/locate/jes](http://www.elsevier.com/locate/jes)

# Effect of relative humidity and the presence of aerosol particles on the $\alpha$ -pinene ozonolysis

Guoqiang Zhang, Hongbo Fu, Jianmin Chen\*

Shanghai Key Laboratory of Atmospheric Particle Pollution and Prevention (LAP<sup>3</sup>), Department of Environmental Science and Engineering, Institute of Atmospheric Sciences, Fudan University, Shanghai 200433, China

## ARTICLE INFO

### Article history:

Received 30 July 2017

Revised 12 October 2017

Accepted 24 October 2017

Available online 8 November 2017

### Keywords:

$\alpha$ -Pinene

Ozonolysis

RH%

Seed aerosol

Rate constant

## ABSTRACT

The  $\alpha$ -pinene ozonolysis under the different environmental conditions were observed in a smog chamber. The second-order rate constant ( $k$ ) was determined to be  $(7.25 \pm 0.06) \times 10^{-17} \text{ cm}^3/(\text{molecule}\cdot\text{sec})$  under 20% of relative humidity (RH) and room temperature. RH showed a marked influence on the  $\alpha$ -pinene ozonolysis. The value of  $k$  increased with RH increase, which was 1.6 times faster at RH = 80% than that at RH = 20%. Additionally, the value of  $k$  apparently changed in the presence of the aerosol particles. The diesel soot increased the  $k$  value. The fly ash prohibited the reaction, however,  $\text{H}_2\text{SO}_4$ -treated fly ash promoted the reaction. The information of products gained using FT-IR and SPAMS showed that pinonic acid, 10-hydroxy-pinonic acid and pinic acid could be generated during the  $\alpha$ -pinene ozonolysis. Water molecules could take part in the formation of the products, and play a vital role in the degradation of  $\alpha$ -pinene. The atmospheric residence time calculation showed that the ozonolysis in the atmosphere is an important way of the  $\alpha$ -pinene consumption as compared to that reacted with OH during daytime. The results suggested that the degradation of  $\alpha$ -pinene via the ozonization in the atmosphere may be affected greatly by RH, as well as the presence of aerosol particles. The ozonolysis reaction may be an important way of the  $\alpha$ -pinene consumption during daytime.

© 2017 The Research Center for Eco-Environmental Sciences, Chinese Academy of Sciences.

Published by Elsevier B.V.

## Introduction

A variety of volatile organic compounds (VOCs) are emitted into the atmosphere from natural sources at a global scale, and the annual global VOCs flux was estimated to be  $11.50 \times 10^{14} \text{ g}$  (Guenther et al., 1995). Among them, the annual global amount of reactive hydrocarbon emissions is  $1.75 \times 10^{14} \text{ g}$  from tree foliage sources, which are much higher by 6.2 times than those from man-made sources (Rasmussen, 1972). Various monoterpenes have remarkable effects in the chemistry of the lower troposphere and the atmospheric boundary layer, their annual natural contribution is  $1.27 \times 10^{14} \text{ g}$  (Guenther et al., 1995). Among monoterpenes emitted from coniferous forests,  $\alpha$ -pinene is

the most important organic compound with the highest emission at an annual rate of  $\sim 0.66 \times 10^{14} \text{ g/year}$  (Guenther et al., 2012). It is well known that  $\alpha$ -pinene could react rapidly with oxidants in the atmosphere (such as OH radicals,  $\text{NO}_3$  radicals and  $\text{O}_3$ ) to produce secondary organic aerosols (SOA) via the gas-phase reactions (Berkemeier et al., 2016; Nah et al., 2016; Rindelaub et al., 2016). Generally, OH radical is a very important photochemical oxidant during daylight hours for the degradation of  $\alpha$ -pinene in the atmosphere, and  $\text{NO}_3$  radical is the major during nighttime hours, whereas  $\text{O}_3$  plays an important role on the  $\alpha$ -pinene ozonolysis throughout the whole day. The  $\alpha$ -pinene ozonolysis could lead to the formation of OH radical, and the measured OH yield was

\* Corresponding author. E-mail: [jmchen@fudan.edu.cn](mailto:jmchen@fudan.edu.cn) (Jianmin Chen).

$0.70 \pm 0.17$ , which was considered to be a non-negligible source of OH (Paulson et al., 1998). The  $\alpha$ -pinene ozonolysis produces many low-volatility compounds that could partition into particle phase and contribute prominently to the formation of SOA (Denjean et al., 2015; Gordon et al., 2015; Zhang et al., 2015).

Relative humidity (RH) is an important factor to influence the formation of SOA produced by the  $\alpha$ -pinene ozonolysis. It was well documented that there was an inverse relationship between the equilibrium partitioning co-efficient and the activity co-efficient for a semivolatile organic compound (Pankow, 1994). Because ambient humidity could have an important influence on the activity co-efficient, the equilibrium partitioning co-efficient on many kinds of aerosols could be affected by RH (Jang and Kamens, 1999; Pankow 1994). Since atmospheric aerosols could absorb water molecules from the gas phase (Vartiainen et al., 1994), the total aerosol yield increases in the case of higher RH due to uptake of water for the experiments of the  $\alpha$ -pinene ozonolysis (Cocker et al., 2001). In addition, the change of RH could affect the mass distribution of the products yielded during the  $\alpha$ -pinene ozonolysis (Fick et al., 2003).

Also, the presence of aerosol shows a marked effect on the formation of SOA from the  $\alpha$ -pinene ozonolysis. Song et al. (2011) reported citric acid significantly enhances the absorption of the oxidized products of  $\alpha$ -pinene into particle-phase, which suggested that a well-mixed organic aerosol may be formed during the  $\alpha$ -pinene ozonolysis. Cocker et al. (2001) found that aqueous  $(\text{NH}_4)_2\text{SO}_4$  aerosols prohibit the SOA formation, whereas Czoschke and Jang (2006) confirmed that the SOA formation are enhanced apparently once the exposure to  $\text{NH}_4\text{HSO}_4$  aerosols in the case of the  $\alpha$ -pinene ozonolysis. The presence of  $(\text{NH}_4)_2\text{SO}_4$  in organic aerosol may affect or inhibit the formation of hemiacetals and acetals, and thus cause a reduction in the aerosol yield.  $\text{NH}_4\text{HSO}_4$  aerosols could initiate the acid catalyzed reactions forming other species (such as hemiacetals and acetals), so could have an increasing effect on the SOA formation. These studies show that different types of aerosol particles exerted the different effects on the SOA formation. Hence, a systematic research concerning the effects of aerosol particles on the SOA formation is essential. Fly ash and diesel soot are ubiquitous in the atmosphere, their effects on the SOA formation are not reported yet, and thus should be studied systematically.

The accurate measurement of the rate constants for the  $\alpha$ -pinene ozonolysis could be used to assess its atmospheric lifetimes, as well as the relative importance of its removal pathways via gas-phase. The rate constants for the  $\alpha$ -pinene ozonolysis had been reported at the different environments (Atkinson et al., 1982; Gordon et al., 2015; Japar et al., 1974; Khamaganov and Hites, 2001; Munshi et al., 1989; Ripperton et al., 1972; Stewart et al., 2013; Witter et al., 2002). Ripperton et al. (1972) early reported that the rate constant of the  $\alpha$ -pinene ozonolysis in 150-L Teflon bag was  $1.70 \times 10^{-18} \text{ cm}^3/\text{cm}^3/(\text{molecule}\cdot\text{sec})$ , while Japar et al. (1974) later determined the value to be  $(3.3 \pm 0.3) \times 10^{-16} \text{ cm}^3/(\text{molecule}\cdot\text{sec})$  using an evacuated 45-L bell jar reactor. Witter et al. (2002) further investigated the kinetic of the  $\alpha$ -pinene ozonolysis under flow-tube conditions, gained the rate constant to be  $(1.1 \pm 0.2) \times 10^{-16} \text{ cm}^3/(\text{molecule}\cdot\text{sec})$ . Stewart et al. (2013) measured rate constants for

the  $\alpha$ -pinene ozonolysis in a static reaction chamber at room temperature, measured the value to be  $(1.07 \pm 0.13) \times 10^{-16} \text{ cm}^3/(\text{molecule}\cdot\text{sec})$ . These results show that remarkable discrepancy exists in the reported rate constants for the  $\alpha$ -pinene ozonolysis under the different experimental conditions. Accordingly, the rate constant for the  $\alpha$ -pinene ozonolysis under the different RH and aerosol particles should be studied systematically.

In this study, the  $\alpha$ -pinene ozonolysis experiments were investigated in an indoor smog chamber. The rate constants of the reactions were determined individually at the different RH (20%, 50% and 80%). The influences of aerosol particles, including fly ash and diesel soot, on the rate constant of the  $\alpha$ -pinene ozonolysis were further observed. On the basis of the product information gained by Fourier Transform infrared spectroscopy (FT-IR) and single particle aerosol mass spectrometry (SPAMS), the reaction mechanism for the  $\alpha$ -pinene ozonolysis was proposed. The effects of RH and aerosols on relative atmospheric residence time of  $\alpha$ -pinene were also calculated.

## 1. Materials and methods

### 1.1. Materials and reagents

Unless otherwise stated, all chemicals were purchased and used without further purification.  $\alpha$ -Pinene ( $\geq 99.0\%$ ) was purchased from the Aladdin. Ethanol was gained from the Sinpharm chemical reagent Co. Ltd. Ozone was formed by ultrapure pure oxygen ( $\geq 99.999\%$ , Shanghai Lou Yang Cylinder Gas Co., Ltd., Shanghai, China) through an ozone generator (Model WFG-8, Shanghai Weizhen Environmental Protection Equipment Factory, Shanghai, China). Water was produced by Milli-Q<sup>®</sup> direct water purification system (Millipore, Bedford, USA).

The fly ash and diesel soot were used in the experiments. The fly ash was attained from rural chimney. The fly ash treated with 0.1 mol/L  $\text{H}_2\text{SO}_4$  was called  $\text{H}_2\text{SO}_4$ -treated fly ash in this study. The absorption peaks at 1140 and 981  $\text{cm}^{-1}$  are observed, indicating that  $\text{SO}_4^{2-}$  ions exist in  $\text{H}_2\text{SO}_4$ -treated fly ash (Givan et al., 1998; Kumar and Kanti, 2007). Diesel soot was collected from the smoke that was produced by burning diesel. The three kinds of aerosols were characterized by using a Nicolet Avatar 360 FT-IR spectrometer equipped with a MCT detector and a diffuse reflectance accessory (see SI Text S1 and Fig. S1 of the Supporting Information).

### 1.2. Chamber experiments

Experiments were carried out using a stainless-steel aerosol chamber equipped with a set of analytical instruments at the room temperature, which has been described in detail elsewhere (Li et al., 2015; Zhang et al., 2011; Li et al., 2017). The volume of aerosol chamber was 4.5  $\text{m}^3$  with a 3.4  $\text{m}^{-1}$  surface-to-volume ratio. The interior wall of aerosol chamber was plated by a 0.3-mm Teflon inner coat. Temperature and RH inside the chamber were measured by a hygroclip monitor (model IM-4, Rotronic). A fan was installed in the bottom of the chamber to ensure the air mixed well during the experiment.

Prior to each test, the chamber was repeatedly evacuated by two vacuum pumps (Frinklin Electric, Model 1,201,006,416, USA), and restored the pressure equilibrium by the HEPA-filtrated

air. After the chamber was scrubbed using aqueous ethanol (50%, v/v), the chamber was evacuated and restored normal pressure once again for reducing RH. Finally the chamber was cleaned using high concentration of O<sub>3</sub> (3 ppm) for 12 hr.

After cleaning, the chamber was evacuated, and was restored the pressure equilibrium with ultrapure air (Shanghai Special Gas Co., Ltd., Shanghai, China). RH in the chamber was controlled by a humidifier, which was consisted of a three-necked flask half filled with Milli-Q grade water (Millipore, Bedford, USA) in the water bath pot at 80°C. Milli-Q grade water vaporized into ultrapure air stream, and then the vapors passed into chamber through a Teflon transfer line. About 2000 molecule/cm<sup>3</sup> polydisperse aerosol particles were then introduced by directly atomizing solid of the particles in the solid aerosol generator (SAG 410, Topas GmbH, Germany) into the chamber. About 600 ppb of O<sub>3</sub> in the chamber was generated by a flow (1 L/min) of ultrapure pure oxygen through an ozone generator and into the chamber. Finally, microliter syringe was used to measure and inject known volumes of liquid  $\alpha$ -pinene (Aladdin,  $\geq 99.0\%$ ) into a thermodenuder (TD) system which was kept at 180°C and vaporized  $\alpha$ -pinene. The gas was then flushed with ultrapure air, and then passed through a Teflon line and into the chamber. The concentration of  $\alpha$ -pinene was approximately 289 ppb in the chamber. All of the experiments were carried out in the dark. Additionally, the blank experiments of ozone and  $\alpha$ -pinene had been done (see SI Text 2-3 and Fig. S2–S4 of the Supporting Information).

### 1.3. Analysis method

The  $\alpha$ -pinene concentration was quantitatively monitored by gas chromatograph-flame ionization detection (GC-FID, GC-9560, Shanghai Huaai Chromatography Analysis Co., Ltd., Shanghai, China) using a 30 m-long, 0.25 mm-i.d. SE-30 dimethylpolysiloxano capillary with a 0.25  $\mu\text{m}$  film thickness. Samples were obtained by drawing 10 L of air through a stainless steel tube containing Tenax TA solid adsorbent. The tube was placed in the inlet of GC, and then heated to 250°C. The column temperature is maintained at 60°C for 5 min, and then ramped at 30°C/min to 270°C, finally was held at 270°C for 5 min. The concentration of  $\alpha$ -pinene was determined from the FID response using a standard curve. Because the concentration of O<sub>3</sub> was much higher than that of  $\alpha$ -pinene, thus  $\alpha$ -pinene was almost completely consumed by O<sub>3</sub> finally. The O<sub>3</sub> concentration in the chamber was continuously monitored by an O<sub>3</sub> analyzer (Thermo Scientific Model 49i, USA).

The SOA formed from the  $\alpha$ -pinene ozonolysis was sampled on the KBr particle by sampling apparatus. Infrared spectra of SOA were recorded in the spectral range from 4000 to 650 cm<sup>-1</sup> using a Nicolet Avatar 360 FT-IR spectrometer which was equipped with a MCT detector cooled by liquid N<sub>2</sub> and a diffuse reflectance accessory. IR spectra were recorded at a resolution of 4 cm<sup>-1</sup>. A 14 mg ( $\pm 0.02$  mg) KBr sample was placed into a ceramic crucible in the chamber.

SPAMS (Hexin Analytical Instrument Co., Ltd., Guangzhou, China) was used in this study, and the technical details of SPAMS had been reported by Li et al. (2011). Briefly, aerosols in the size range of 0.2–2.0  $\mu\text{m}$  were drawn into the aerodynamic lens through a 0.1 mm critical orifice at a flow of 80 mL/min, and then focused into a collimated tiny particle beam. Then

particles attained a terminal size-dependent aerodynamic velocity, which was measured by detection of light scattered from two continuous lasers (532 nm) separated by a fixed 6 cm distance in the sizing chamber. The particle arriving at the ion source region was desorbed and ionized by a pulsed laser (Q-switched Nd:YAG laser, 266 nm). The positive and negative ions of the particle were detected simultaneously by the time-of-flight mass spectrometer. Particle diameter was calculated from the measured speed and calibrated with polystyrene latex spheres (PSLs) with known diameters.

### 1.4. Rate constant measurements

Rate constants for the  $\alpha$ -pinene ozonolysis under the different conditions were determined by monitoring the concentrations of  $\alpha$ -pinene and O<sub>3</sub>. The removal processes of O<sub>3</sub> and  $\alpha$ -pinene under different conditions were following Reaction 1.



The rate equation for Reaction 1 is given by,

$$\frac{dx}{dt} = k(a-x)(b-x) \quad (2)$$

where  $k$  is the rate constant,  $a$  and  $b$  are the initial concentrations of  $\alpha$ -pinene and O<sub>3</sub>, and  $x$  is the consumed concentrations at time  $t$ .

When the concentrations of  $\alpha$ -pinene and O<sub>3</sub> are different, that is to say  $a \neq b$ , then the rate Eq. (2) is integrated, the equation was given by

$$\frac{1}{a-b} \ln \frac{b(a-x)}{a(b-x)} = kt \quad (3)$$

Clearly, a plot of  $\ln\{b(a-x)/[a(b-x)]\}$  versus  $t$  will have a slope equal to the value of  $(a-b)k$ .

## 2. Results and discussion

### 2.1. Effect of RH on the rate constant

The dots of  $\ln\{b(a-x)/[a(b-x)]\}$  versus  $t$  under the different RH could be drawn up, and then the values of  $k$  could be derived through the slopes of the linear least-squares best fits to the data, finally are listed in Table 1. The best-fit lines are not forced through the origin. The second-order rate constants are  $(7.25 \pm 0.06) \times 10^{-17}$  at RH = 20%,  $(9.78 \pm 0.59) \times 10^{-17}$  at RH = 50%, and  $(1.18 \pm 0.06) \times 10^{-16}$  at RH = 80% in units of cm<sup>3</sup>/(molecule-sec), respectively. Under this condition, one can see that the degradation of  $\alpha$ -pinene enhances with increasing RH.

It was well known that RH could affect the mass distribution of product formed by the  $\alpha$ -pinene ozonolysis (Fick et al., 2003). Furthermore, RH plays an important role on the mechanism of the  $\alpha$ -pinene ozonolysis, and could impact the rate constant  $k$  value of the  $\alpha$ -pinene ozonolysis. The increasing of RH value could decrease viscosity of the amorphous aerosol particles, and then increase the oxidant uptake due to the enhanced particle-phase diffusion (Chan and Chan,

**Table 1 – Rate constants for the reaction between  $\alpha$ -pinene and  $O_3$  under different RH.**

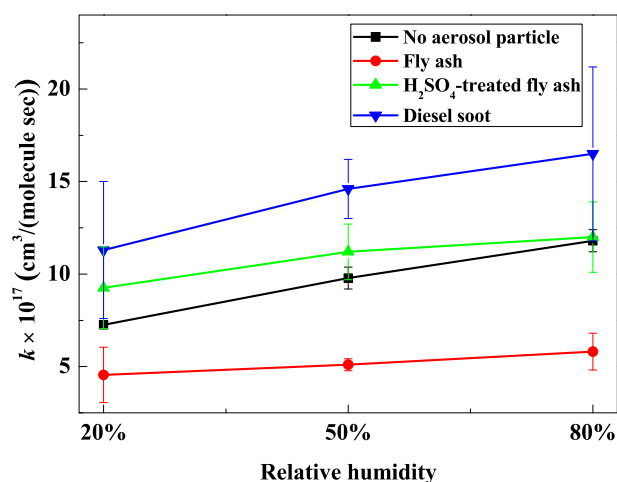
| Reaction conditions | $k$ ( $cm^3/(\text{molecule}\cdot\text{sec})$ ) |
|---------------------|---|
| RH = 20%            | $(7.25 \pm 0.06) \times 10^{-17}$               |
| RH = 50%            | $(9.78 \pm 0.59) \times 10^{-17}$               |
| RH = 80%            | $(1.18 \pm 0.06) \times 10^{-16}$               |

RH: relative humidity.

2012; Kuwata and Martin, 2012; Liu et al., 2014). On the other hand, some researchers observed that increasing RH could prohibit the oxidant uptake due to the competition between water and oxidant uptake, because the reactive sites were taken up by water molecules (Slade and Knopf, 2014). For this study, the viscosity of the seed aerosol particles could be lowered with increasing RH, subsequently the diffusion of oxidant on particle-phase could be enhanced. Finally, the  $O_3$  uptake could be promoted, and the rate constant of  $k$  value for the  $\alpha$ -pinene ozonolysis increases with RH.

## 2.2. Effect of aerosol particles on the rate constant

The rate constant as a function of RH value in the presence of aerosol particles are shown in Fig. 1. When RH value is set at 20%, the degradation of  $\alpha$ -pinene is promoted by diesel soot, and prohibited by fly ash. However, the  $\alpha$ -pinene degradation is promoted by the  $H_2SO_4$ -treated fly ash. The rate constant is  $(1.13 \pm 0.37) \times 10^{-16} \text{ cm}^3/(\text{molecule}\cdot\text{sec})$  in the case of diesel soot, which is much high than that in the absence of the aerosols  $((7.25 \pm 0.06) \times 10^{-17} \text{ cm}^3/(\text{molecule}\cdot\text{sec}))$ . The rate constant changes to  $(4.55 \pm 1.50) \times 10^{-17} \text{ cm}^3/(\text{molecule}\cdot\text{sec})$ , when fly ash is used. This suggests that the presence of fly ash particles apparently prohibits the  $\alpha$ -pinene ozonolysis. However, the rate constant changes to  $(1.12 \pm 0.15) \times 10^{-16} \text{ cm}^3/(\text{molecule}\cdot\text{sec})$ , when  $H_2SO_4$ -treated fly ash is used. When RH value is set at 50% and 80%, respectively, diesel soot could promote the  $\alpha$ -pinene ozonolysis, and fly ash prohibits it. However,  $H_2SO_4$ -treated fly ash could promote the  $\alpha$ -pinene ozonolysis.



**Fig. 1 – Representative curves for rate constant of the  $\alpha$ -pinene ozonolysis with different aerosol particles as a function of RH value. RH: relative humidity.**

For some solid catalysts, water molecules sometimes are detrimental, due to occupying reactive sites, and thus decrease catalytic activities (Okuhara, 2002). Diesel soot is hydrophobic, so its catalytic activity may be affected slightly, and is the highest. Hence, in the case of the diesel soot, the second-order rate constants are the biggest. Fly ash is hydrophilic, so the catalytic activity is suppressed obviously and is the lowest. As for the experiments with fly ash, the second-order rate constants are thus the smallest one. The  $H_2SO_4$ -treated fly ash is hydrophilic, so the catalytic activity is suppressed obviously. But they are treated by  $H_2SO_4$  solution, so they could have more reactive sites, the catalytic activity is enhanced partly. Hence, in the presence of the  $H_2SO_4$ -treated fly ash, the second-order rate constants are bigger than that without aerosol particle.

Rate constants for the  $\alpha$ -pinene ozonolysis available from the literature are compared with those from the study in Table 2. The data shown herein are in excellent agreement with those of Stewart et al. (2013), Witter et al. (2002), Atkinson et al. (1982 and 1990), Munshi et al. (1989), and Nolting et al. (1988). However, remarkable discrepancy exists in the reported rate constants from Japar et al. (1974) and Ripperton et al. (1972) and the data shown herein. The discrepancy is probably caused by systematic experimental factors.

## 2.3. Product identifications for the $\alpha$ -pinene ozonolysis

FT-IR spectra of products from the  $\alpha$ -pinene ozonolysis at room temperature and RH = 50% are shown in Fig. 2. The functional groups corresponding to characteristic absorption peaks are assigned based upon the previous researches (Ataol and Ekici, 2014; Doroshenko et al., 2010; Gómez-Serrano et al., 1999; Ibrahim et al., 2005; Maroń et al., 2011). These infrared spectra provide valuable information on different surface species adsorbed on the KBr particles by the vibrational modes.

**Table 2 – Rate constants for the  $\alpha$ -pinene ozonolysis in literature.**

| Reaction conditions   | $k$ ( $cm^3/(\text{molecule}\cdot\text{sec})$ ) | Reference               |
|---|---|-------------------------|
| $C_{O_3} = 1\text{--}6$ ppm,<br>$C_{\alpha\text{-pinene}} = 10$ ppm,<br>$T = 25 \pm 2^\circ\text{C}$ , RH < 5%                  | $(1.07 \pm 0.13) \times 10^{-16}$               | Stewart et al. (2013)   |
| $C_{O_3} = 0.122\text{--}1.88$ ppm,<br>$C_{\alpha\text{-pinene}} = 78\text{--}67$ ppb,<br>$T = 23^\circ\text{C}$ , RH < 5%      | $(1.1 \pm 0.2) \times 10^{-16}$                 | Witter et al. (2002)    |
| $C_{\alpha\text{-pinene}} \gg C_{O_3}$<br>$T = 23^\circ\text{C}$ , RH < 1%  | $8.7 \times 10^{-17}$                           | Atkinson et al. (1990)  |
| $C_{O_3} = 1391$ ppm,<br>$C_{\alpha\text{-pinene}} = 290$ ppm,<br>$T = 25^\circ\text{C}$ , RH < 5%                              | $8.0 \times 10^{-17}$                           | Munshi et al. (1989)    |
| $C_{O_3} = 20\text{--}50$ ppm,<br>$C_{\alpha\text{-pinene}} = 10\text{--}200$ ppb<br>$T = 28^\circ\text{C}$ , RH < 5%           | $(8.6 \pm 1.3) \times 10^{-17}$                 | Nolting et al. (1988)   |
| $C_{O_3} = 0.2\text{--}0.4$ ppm,<br>$C_{\alpha\text{-pinene}} = 2\text{--}4$ ppm,<br>$T = 3\text{--}48^\circ\text{C}$ , RH < 1% | $8.4 \times 10^{-17}$                           | Atkinson et al. (1982)  |
| $C_{O_3} = 25\text{--}60$ ppb,<br>$C_{\alpha\text{-pinene}} = 214\text{--}537$ ppm,<br>$T = 25^\circ\text{C}$ , RH < 5%         | $(3.3 \pm 0.3) \times 10^{-16}$                 | Japar et al. (1974)     |
| $C_{O_3} = 4.1\text{--}26.5$ ppm,<br>$C_{\alpha\text{-pinene}} = 64.9\text{--}121.1$ ppm,<br>$T = 21^\circ\text{C}$ , RH < 5%   | $1.70 \times 10^{-18}$                          | Ripperton et al. (1972) |

The absorption peaks at 3427, 3438, and 3477  $\text{cm}^{-1}$  are observed, corresponding to the O–H stretching vibration in alcohol (Doroshenko et al., 2010). The absorption peaks at 2849, 2850, 2913, and 2917  $\text{cm}^{-1}$  are assigned as O–H stretching vibration in carboxylic acid (Maroñ et al., 2011). The absorption peaks appear at 1649, 1660, 1741, and 1743  $\text{cm}^{-1}$ , due to C=O stretching vibration (Ibrahim et al., 2005; Lee and RAC, 1999). The absorption peaks at 1384, 1385, and 1390  $\text{cm}^{-1}$  are attributed to  $\text{CH}_3$  the bending vibration in alkane (Gómez-Serrano et al., 1999). The spectral characteristic of C–H in the cyclobutyl appears at about 1093  $\text{cm}^{-1}$  (Ataol and Ekici, 2014). The absorption peaks at about 930 and 914  $\text{cm}^{-1}$  are assigned to the O–H deformation vibration for carboxylic acid (Ibrahim et al., 2005).

Also, SPAMS was used to identify the possible products from the  $\alpha$ -pinene ozonolysis. The average negative mass spectrum of products from the  $\alpha$ -pinene ozonolysis under 50% of RH and room temperature is shown in Fig. 3. The characteristic  $m/z$  and possible products are assigned based upon the previous studies (Camredon et al., 2010; Claeys et al., 2009; Peng et al., 2016; Jaoui and Kamens, 2001), and are shown in Table S1. In Fig. 3 average negative mass spectrum at  $m/z = -183$ ,  $-185$ ,  $-189$ ,  $-199$  and  $-213$  could be ascribed to  $\text{C}_{10}\text{H}_{16}\text{O}_3$  (Camredon et al., 2010; Jaoui and Kamens, 2001; Jenkin et al., 2000; Kristensen et al., 2013; Ma et al., 2008; Witkowski and Gierczak, 2014),  $\text{C}_9\text{H}_{14}\text{O}_4$  (Camredon et al., 2010; Jaoui and Kamens, 2001; Jenkin et al., 2000; Kristensen et al., 2013; Ma et al., 2008; Witkowski and Gierczak, 2014),  $\text{C}_8\text{H}_{14}\text{O}_5$  (Claeys et al., 2009;

Kristensen et al., 2013; Witkowski and Gierczak, 2014; Yasmeen et al., 2010),  $\text{C}_{10}\text{H}_{16}\text{O}_4$  (Camredon et al., 2010; Jaoui and Kamens, 2001; Jenkin et al., 2000; Kristensen et al., 2013; Ma et al., 2008; Witkowski and Gierczak, 2014) and  $\text{C}_{10}\text{H}_{14}\text{O}_5$  (Peng et al. 2016), respectively. The functional groups in these products include OH, C=O,  $\text{CH}_3$ , cyclobutyl groups and so on. These results gained by SPAMS are in excellent agreement with those attained by FT-IR.

Fig. 4 presents the proposed mechanism of the  $\alpha$ -pinene ozonolysis (Docherty et al., 2005; Jang and Kamens, 1999; Jenkin, 2004; Jenkin et al., 2000; Kurtén et al., 2015). Energy-rich ozonides are generated from the addition of  $\text{O}_3$  to the double bond of  $\alpha$ -pinene initially, and then decompose rapidly by two possible channels to form two different carbonyl-substituted Criegee intermediates of CI-1 and CI-2. CI-1 reacts with water initially, and then loses water, finally forms pinonic acid (route-1). The mechanism producing the pinonic acid shows that water takes part in the process, suggesting that water could have an important effect on the rate constant for the  $\alpha$ -pinene ozonolysis. It was well documented that Criegee intermediate may react primarily with water vapor under atmospheric conditions (Neeb et al., 1996; Tobias and Ziemann, 2001). A theoretical study showed that the observed rate constant for the reaction of Criegee is linearly dependent on the water concentration, which increases with RH increasing (Ryzhkov and Ariya, 2004). High RH was thus lead to the increase of the rate constant for the  $\alpha$ -pinene ozonolysis.

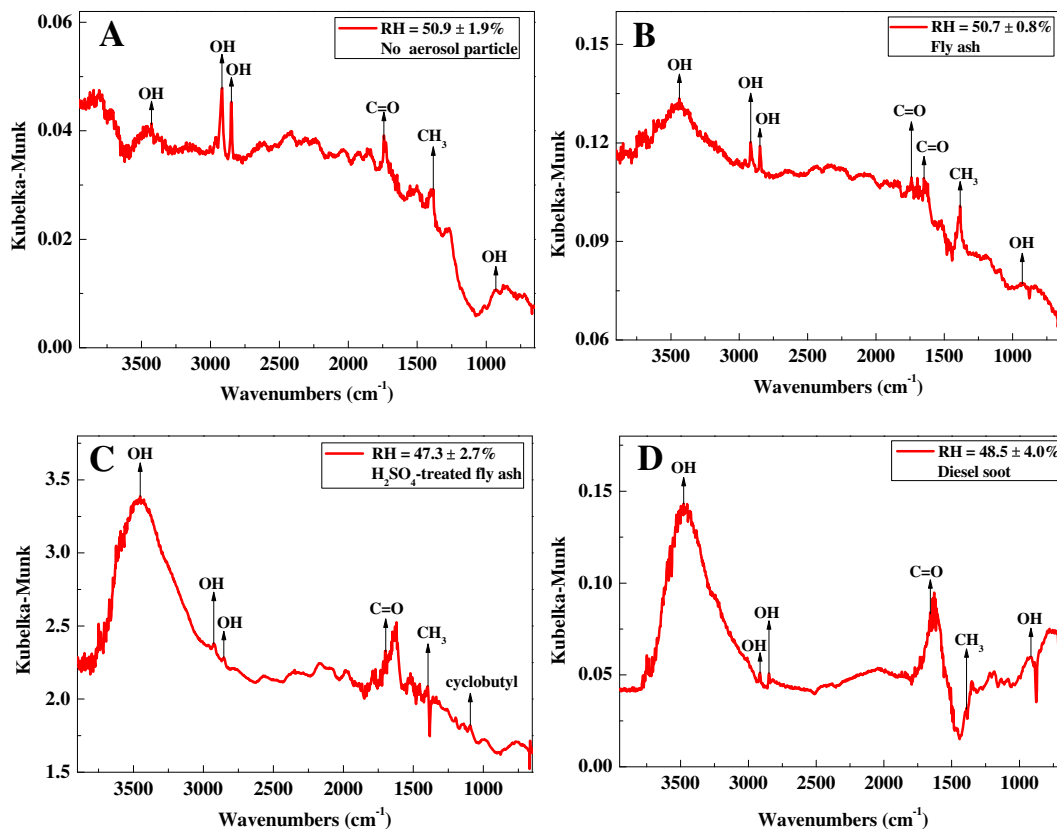
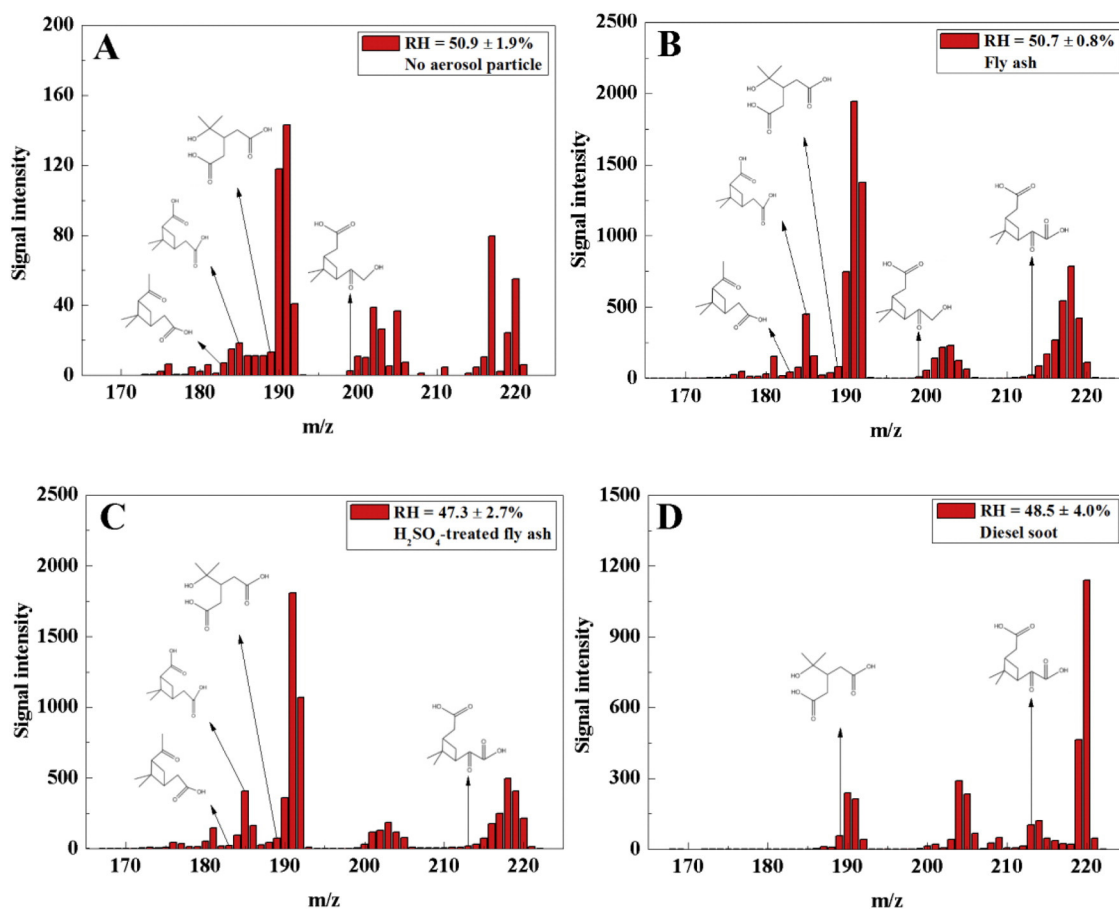


Fig. 2 – Fourier Transform infrared spectroscopy (FR-IR) spectra of the products formed by the  $\alpha$ -pinene ozonolysis in different conditions: (A) in the absence of aerosol particle; (B) in the presence of fly ash; (C) in the presence of  $\text{H}_2\text{SO}_4$ -treated fly ash; (D) in the presence of diesel soot.



**Fig. 3** – Average negative mass spectrum for SOA formed from the  $\alpha$ -pinene ozonolysis under different experimental conditions: (A) in the absence of aerosol particle; (B) in the presence of fly ash; (C) in the presence of  $\text{H}_2\text{SO}_4$ -treated fly ash; (D) in the presence of diesel soot.

1,4H shifts of CI-2 generate a vinylhydroperoxide, which dissociates to produce a OH radical and a vinoxy radical. The addition of  $\text{O}_2$  to the vinoxy radical leads to a peroxy radical. The reaction of the peroxy radical with  $\text{RO}_2$  produces a  $\text{C}_{10}$  alkoxy radical, which forms the  $\text{C}_{10}$  acyl radical via the isomerisation reaction of 1,8H atom shift. The addition of  $\text{O}_2$  to the  $\text{C}_{10}$  acyl radical leads to a  $\text{C}_{10}$  acyl peroxy radical, which reacts with the  $\text{RO}_2/\text{HO}_2$  to generate a  $\text{C}_{10}$ -hydroxypinic acid (route-2).

The thermal decomposition of the  $\text{C}_{10}$  alkoxy radical produces a formaldehyde (HCHO) and a  $\text{C}_9$  acyl radical. The addition of  $\text{O}_2$  to the  $\text{C}_9$  acyl radical produces a  $\text{C}_9$  acyl peroxy radical, which reacts with the  $\text{RO}_2$  to generate a  $\text{C}_9$  acyloxy radical. The acyloxy radical transforms into a  $\text{C}_9$  acyl radical via isomerization by a 1,7H shift. The addition of  $\text{O}_2$  to the  $\text{C}_9$  acyl radical forms the  $\text{C}_9$  acyl peroxy radical, which reacts with the  $\text{RO}_2/\text{HO}_2$  to produce pinic acid (route-3).

#### 2.4. Relative atmospheric residence times

Relative atmospheric residence time for a kind of VOC is time remaining in the atmosphere. It is relevant with kinds and concentration of oxidants. From the rate constants obtained in this work and references, atmospheric lifetime oxidation by

$\text{O}_3$ , OH and  $\text{NO}_3$  could be calculated for  $\alpha$ -pinene from the following equation:

$$\tau_{\text{oc}} = \frac{1}{k \times [\text{oxidant}]} \quad (4)$$

In the present work, 40 ppb ( $9.84 \times 10^{11}$  molecule/ $\text{cm}^3$ ) was chose as a background value of  $\text{O}_3$  concentration (Singh et al., 1978),  $1 \times 10^6$  molecules/ $\text{cm}^3$  was assumed as OH radical concentration during day-time (Finlayson-Pitts and Pitts, 2000), and 10 ppt ( $2.46 \times 10^8$  molecule/ $\text{cm}^3$ ) was selected as a  $\text{NO}_3$  radical concentration during night-time (Stewart et al., 2013). The lifetimes for  $\alpha$ -pinene are shown in Table 3.

Comparing with the value for the experiment without aerosol particle, atmospheric lifetime of  $\alpha$ -pinene increased in the presence of the fly ash, whereas decreased in the presence of the  $\text{H}_2\text{SO}_4$ -treated fly ash and diesel soot at RH = 20%, 50% and 80%.

Under different conditions, the atmospheric lifetime of  $\alpha$ -pinene oxidized by  $\text{O}_3$  which is dependent on RH and aerosol particles, varies from 1.71 to 6.43 hr. On the basis of the data reported by Stewart et al. (2013) and Hallquist et al. (1999), the atmospheric lifetime of  $\alpha$ -pinene oxidized by OH and  $\text{NO}_3$  radicals are 5.22 and 0.22 hr, respectively. Thus, the

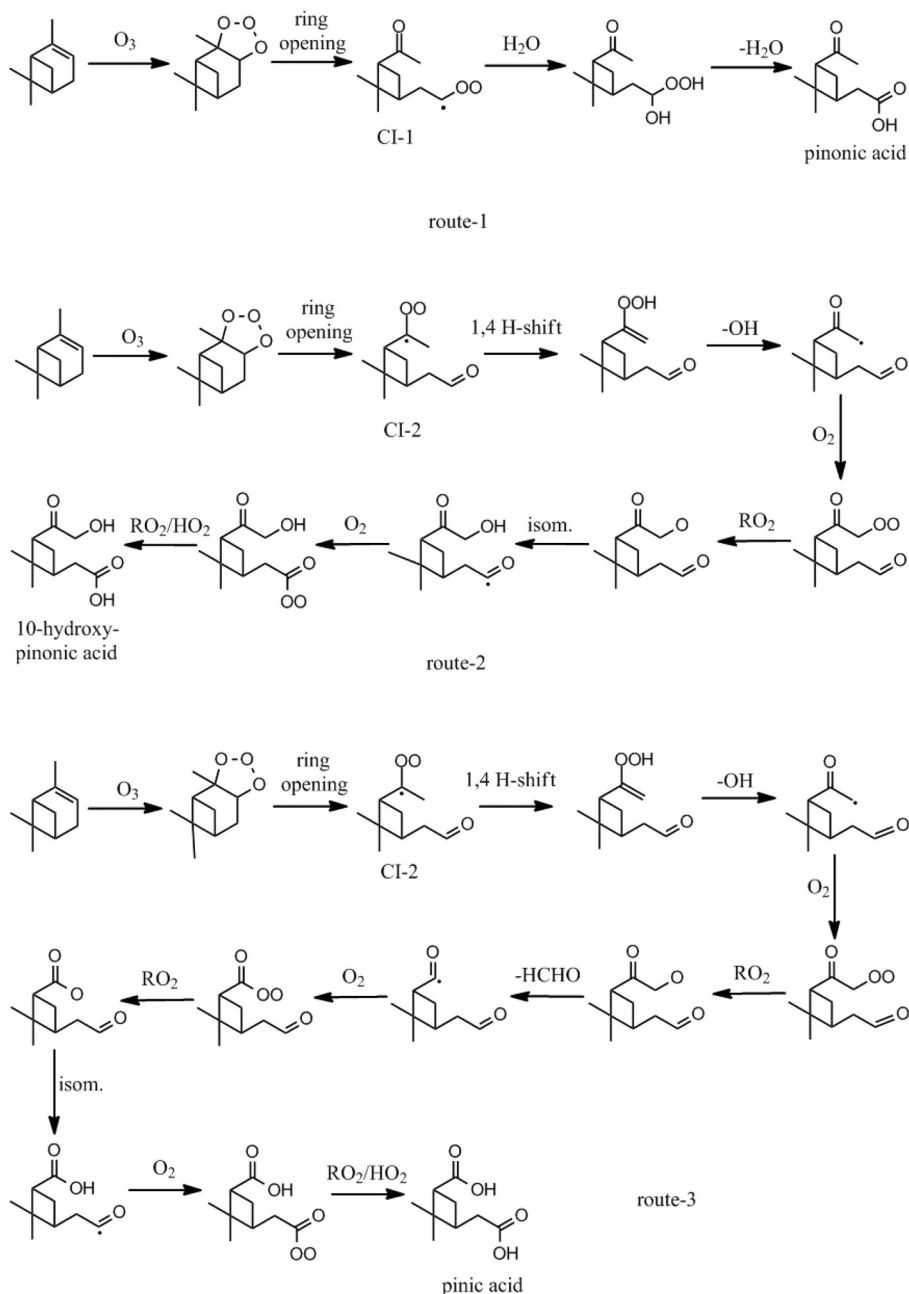


Fig. 4 – The reaction mechanism of the  $\alpha$ -pinene ozonolysis.

ozonolysis in the atmosphere is an important way of the  $\alpha$ -pinene consumption as compared to that via the reaction of  $\alpha$ -pinene and OH during daytime, while  $\alpha$ -pinene is primarily oxidized by the  $NO_3$  radical at night.

### 3. Conclusions

In this study, the  $\alpha$ -pinene ozonolysis had been investigated using a chamber at the different experimental conditions. RH and aerosol particles both have important effects on the rate constant for the  $\alpha$ -pinene ozonolysis under the investigated conditions. The second-order rate constant ( $k$ ) of the  $\alpha$ -pinene ozonolysis increases with the increasing of RH.  $k$  of the  $\alpha$ -

pinene ozonolysis increased in the presence of diesel soot and the  $H_2SO_4$ -treated fly ash, however, the value decreased in the case of the untreated fly ash. FT-IR results show that the  $\alpha$ -pinene ozonolysis could form the products with OH, C=O and COOH groups, leading to poly(carboxylic acid)s. SPAMS data show that pinonic acid, pinic acid, diaterpenylic acid, OH-pinonic acid, and  $C_{10}H_{14}O_5$  may be involved in the  $\alpha$ -pinene ozonolysis. The proposed reaction mechanism for the  $\alpha$ -pinene ozonolysis shows water molecules may play an important role in the degradation rate of  $\alpha$ -pinene. Calculated atmospheric lifetimes suggest the  $\alpha$ -pinene ozonolysis to be an important loss way for  $\alpha$ -pinene during daytime.

This study has very significant atmospheric implications. The concentration measured in the field measurements and

**Table 3 – Calculated atmospheric lifetimes of  $\alpha$ -pinene with respect to oxidation by  $O_3$ ,  $NO_3$  and OH radicals.**

| Rate constants (cm <sup>3</sup> /(molecule·sec))   | Oxidant concentration (molecules/cm <sup>3</sup> ) | Atmospheric lifetimes of $\alpha$ -pinene (hr) |
|--|--|--|
| No aerosol particle<br>$k = (0.725, 0.978, 1.18) \times 10^{-16}$  | $9.84 \times 10^{11}$                              | 3.89, 2.89, 2.39                               |
| Fly ash<br>$k = (4.55, 5.1, 5.81) \times 10^{-17}$   | $9.84 \times 10^{11}$                              | 6.20, 5.54, 4.89                               |
| H <sub>2</sub> SO <sub>4</sub> -treated fly ash<br>$k = (0.926, 1.12, 1.20) \times 10^{-16}$                         | $9.84 \times 10^{11}$                              | 3.05, 2.52, 2.35                               |
| Diesel soot<br>$k = (1.13, 1.46, 1.65) \times 10^{-16}$<br>$k_{OH} = 5.32 \times 10^{-11}$<br>(Stewart et al., 2013) | $9.84 \times 10^{11}$<br>$1 \times 10^6$           | 2.5, 1.93, 1.71<br>5.22                        |

the relative atmospheric residence time for  $\alpha$ -pinene may be affected by RH and aerosol particles in the troposphere.

## Acknowledgments

This work was supported by the Ministry of Science and Technology of China (Nos. 2016YFC0202700, 2016YFE0112200, 2014BAC22B01), the National Natural Science Foundation of China (Nos. 21527814, 21077025) and the Marie Skłodowska-Curie Actions (Nos. 690958-MARSU-RISE-2015).

## Appendix A. Supplementary data

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.jes.2017.10.011>.

## REFERENCES

- Ataol, Ç.Y., Ekici, Ö., 2014. Experimental and theoretical studies of (FTIR, FT-NMR, UV-visible, X-ray and DFT) 2-(4-Allyl-5-pyridin-4-yl-4H-[1,2,4]triazol-3-ylsulfanyl)-1-(3-methyl-3-phenylcyclobutyl)-ethanone. *J. Mol. Struct.* 1065–1066 (1), 1–9.
- Atkinson, R., Winer, A.M., Pitts, J.N., 1982. Rate constants for the gas phase reactions of  $O_3$  with the natural hydrocarbons isoprene and  $\alpha$ - and  $\beta$ -pinene. *Atmos. Environ.* 16 (5), 1017–1020.
- Atkinson, R., Hasegawa, D., Aschmann, S.M., 1990. Rate constants for the gas-phase reactions of  $O_3$  with a series of monoterpenes and related compounds at  $296 \pm 2$  K. *In. J. Chem. Kinet.* 22 (8), 871–887.
- Berkemeier, T., Ammann, M., Mentel, T.F., Pöschl, U., Shiraiwa, M., 2016. Organic nitrate contribution to new particle formation and growth in secondary organic aerosols from  $\alpha$ -pinene ozonolysis. *Environ. Sci. Technol.* 50 (12), 6334–6342.
- Camredon, M., Hamilton, J.F., Alam, M.S., Wyche, K.P., Carr, T., White, I.R., et al., 2010. Distribution of gaseous and particulate organic composition during dark  $\alpha$ -pinene ozonolysis. *Atmos. Chem. Phys.* 10 (6), 2893–2917.
- Chan, L.P., Chan, C.K., 2012. Displacement of ammonium from aerosol particles by uptake of triethylamine. *Aerosol Sci. Tech.* 46 (2), 236–247.
- Claeys, M., Iinuma, Y., Szmigielski, R., Surratt, J.D., Blockhuys, F., Van Alsenoy, C., et al., 2009. Terpenylic acid and related compounds from the oxidation of  $\alpha$ -pinene: implications for new particle formation and growth above forests. *Environ. Sci. Technol.* 43 (18), 6976–6982.
- Cocker, I.I.I.D.R., Clegg, S.L., Flagan, R.C., Seinfeld, J.H., 2001. The effect of water on gas-particle partitioning of secondary organic aerosol. Part I:  $\alpha$ -Pinene/ozone system. *Atmos. Environ.* 35 (12), 6049–6072.
- Czochke, N.M., Jang, M., 2006. Acidity effects on the formation of  $\alpha$ -pinene ozone SOA in the presence of inorganic seed. *Atmos. Environ.* 40 (23), 4370–4380.
- Denjean, C., Formenti, P., Picquet-Varrault, B., Panguì, E., Zapf, P., Katrib, Y., et al., 2015. Relating hygroscopicity and optical properties to chemical composition and structure of secondary organic aerosol particles generated from the ozonolysis of  $\alpha$ -pinene. *Atmos. Chem. Phys.* 15 (6), 3339–3358.
- Docherty, K.S., Wu, W., Lim, Y.B., Ziemann, P.J., 2005. Contributions of organic peroxides to secondary aerosol formed from reactions of monoterpenes with  $O_3$ . *Environ. Sci. Technol.* 39 (11), 4049–4059.
- Doroshenko, I., Pogorelov, V., Sablinskas, V., Balevicius, V., 2010. Matrix-isolation study of cluster formation in methanol: O–H stretching region. *J. Mol. Liq.* 157 (2–3), 142–145.
- Fick, J., Pommer, L., Nilsson, C., Andersson, B., 2003. Effect of OH radicals, relative humidity, and time on the composition of the products formed in the ozonolysis of  $\alpha$ -pinene. *Atmos. Environ.* 37 (29), 4087–4096.
- Finlayson-Pitts, B.J., Pitts, J.N., 2000. *Chemistry of the Upper and Lower Atmosphere*. Academic Press, San Diego.
- Givan, A., Larsen, L.A., Loewenschuss, A., Nielsen, C.J., 1998. Infrared matrix isolation study of H<sub>2</sub>SO<sub>4</sub> and its complexes with H<sub>2</sub>O. *J. Chem. Soc.* 94 (7), 927–935.
- Gómez-Serrano, V., Piriz-Almeida, F., Durán-Valle, C.J., Pastor-Villegas, J., 1999. Formation of oxygen structures by air activation. A study by FT-IR spectroscopy. *Carbon* 37 (10), 1517–1528.
- Gordon, C.A., Ye, J., Chan, A.W.H., 2015. Secondary organic aerosol formation enhanced by organic seeds of similar polarity at atmospherically relative humidity. *SFJ* 1 (2), 6–10.
- Guenther, A., Hewitt, C.N., Erickson, D., Fall, R., Geron, C., Graedel, T., et al., 1995. A global model of natural volatile organic compound emissions. *J. Geophys. Res.* 100 (D5), 8873–8892.
- Guenther, A.B., Jiang, X., Heald, C.L., Sakulyanontvittaya, T., Duhl, T., Emmons, L.K., et al., 2012. The model of emissions of gases and aerosols from nature version 2.1 (MEGAN2.1): an extended and updated framework for modeling biogenic emissions. *Geosci. Model Dev.* 5 (6), 1471–1492.
- Hallquist, M., Wangberg, I., Ljungstrom, E., Barnes, I., Becker, K.H., 1999. Aerosol and product yields from  $NO_3$  radical-initiated oxidation of selected monoterpenes. *Environ. Sci. Technol.* 33 (4), 553–559.
- Ibrahim, M., Nada, A., Kamal, D.E., 2005. Density functional theory and FT-IR spectroscopic study of carboxyl group. *Indian J. Pure Appl. Phys.* 43 (12), 911–917.
- Jang, M., Kamens, R.M., 1999. Newly characterized products and composition of secondary aerosols from the reaction of  $\alpha$ -pinene with ozone. *Atmos. Environ.* 33 (3), 459–474.
- Jaoui, M., Kamens, R.M., 2001. Mass balance of gaseous and particulate products analysis from  $\alpha$ -pinene/ $NO_3$ /air in the presence of natural sunlight. *J. Geophys. Res. Atmos.* 106 (D12), 12541–12558.
- Japar, S.M., Wu, C.H., Niki, H., 1974. Rate constants for the gas phase reaction of ozone with  $\alpha$ -pinene and terpinolene. *Environ. Lett.* 7 (3), 245–249.



- Jenkin, M.E., 2004. Modelling the formation and composition of secondary organic aerosol from  $\alpha$ - and  $\beta$ -pinene ozonolysis using MCM v3. *Atmos. Chem. Phys.* 4 (7), 1741–1757.
- Jenkin, M.E., Shallcross, D.E., Harvey, J.N., 2000. Development and application of a possible mechanism for the generation of cis-pinic acid from the ozonolysis of  $\alpha$ - and  $\beta$ -pinene. *Atmos. Environ.* 34 (18), 2837–2850.
- Khamaganov, V.G., Hites, R.A., 2001. Rate constants for the gas-phase reactions of ozone with isoprene,  $\alpha$ - and  $\beta$ -pinene, and limonene as a function of temperature. *J. Phys. Chem. A* 105 (4), 815–822.
- Kristensen, K., Enggrob, K.L., King, S.M., Worton, D.R., Platt, S.M., Mortensen, R., et al., 2013. Formation and occurrence of dimer esters of pinene oxidation products in atmospheric aerosols. *Atmos. Chem. Phys.* 13 (7), 3763–3776.
- Kumar, V.S., Kanti, D.M., 2007. Direct and rapid determination of sulphate in environmental samples with diffuse reflectance Fourier transform infrared spectroscopy using KBr substrate. *Talanta* 71 (4), 1546–1552.
- Kurtén, T., Rissanen, M.P., Mackeprang, K., Thornton, J.A., Hyttinen, N., Jørgensen, S., et al., 2015. Computational study of hydrogen shifts and ring-opening mechanisms in  $\alpha$ -pinene ozonolysis products. *J. Phys. Chem. A* 119 (46), 11366–11375.
- Kuwata, M., Martin, S.T., 2012. Phase of atmospheric secondary organic material affects its reactivity. *Proc. Natl. Acad. Sci. U. S. A.* 109 (43), 17354–17359.
- Lee, D.H., RAC Sr., 1999. FT-IR spectral characterization of thin film coatings of oleic acid on glasses: I. Coatings on glasses from ethyl alcohol. *J. Mater. Sci.* 34 (1), 139–146.
- Li, L., Huang, Z., Dong, J., Li, M., Gao, W., Nian, H., et al., 2011. Real time bipolar time-of-flight mass spectrometer for analyzing single aerosol particles. *Int. J. Mass Spectrom.* 303 (2), 118–124.
- Li, C., Ma, Z., Chen, J., Wang, X., Ye, X., Wang, L., et al., 2015. Evolution of biomass burning smoke particles in the dark. *Atmos. Environ.* 120, 244–252.
- Li, C., Hu, Y., Zhang, F., Chen, J., Ma, Z., Ye, X., et al., 2017. Multi-pollutant emissions from the burning of major agricultural residues in China and the related health-economic effects. *Atmos. Chem. Phys.* 17 (8), 4957–4988.
- Liu, Y., Huang, L., Li, S.M., Harner, T., Liggio, J., 2014. OH-initiated heterogeneous oxidation of tris-2-butoxyethyl phosphate: implications for its fate in the atmosphere. *Atmos. Chem. Phys.* 14 (22), 12195–12207.
- Ma, Y., Russell, A.T., Marston, G., 2008. Mechanisms for the formation of secondary organic aerosol components from the gas-phase ozonolysis of  $\alpha$ -pinene. *Phys. Chem. Chem. Phys.* 10 (29), 4294–4312.
- Maroń, M.K., Takahashi, K., Shoemaker, R.K., Vaida, V., 2011. Hydration of pyruvic acid to its geminal-diol, 2,2-dihydroxypropanoic acid, in a water-restricted environment. *Chem. Phys. Lett.* 513 (4–6), 184–190.
- Munshi, H.B., Rao, K., Iyer, R.M., 1989. Rate constants of the reactions of ozone with nitriles, acrylates and terpenes in gas-phase. *Atmos. Environ.* 23 (9), 1971–1976.
- Nah, T., Sanchez, J., Boyd, C.M., Ng, N.L., 2016. Photochemical aging of  $\alpha$ -pinene and  $\beta$ -pinene secondary organic aerosol formed from nitrate radical oxidation. *Environ. Sci. Technol.* 50 (1), 222–231.
- Neeb, P., Horie, O., Moortgat, G.K., 1996. Gas-phase ozonolysis of ethene in the presence of hydroxylic compounds. In: *J. Chem. Kinet.* 28 (10), 721–730.
- Nolting, F., Behnke, W., Zetzsch, C., 1988. A smog chamber for studies of the reactions of terpenes and alkanes with ozone and OH. *J. Atmos. Chem.* 6 (1), 47–59.
- Okuhara, T., 2002. Water-tolerant solid acid catalysts. *Chem. Rev.* 102 (10), 3641–3666.
- Pankow, J.F., 1994. An absorption model of gas/particle partitioning of organic compounds in the atmosphere. *Atmos. Environ.* 28 (2), 185–188.
- Paulson, S.E., Chung, M., Sen, A.D., Orzechowska, G., 1998. Measurement of OH radical formation from the reaction of ozone with several biogenic alkenes. *J. Geophys. Res.* 103 (D19), 25533–25539.
- Peng, J., Hu, M., Guo, S., Du, Z., Zheng, J., Shang, D., et al., 2016. Markedly enhanced absorption and direct radiative forcing of black carbon under polluted urban environments. *Proc. Natl. Acad. Sci. U. S. A.* 113 (16), 4266–4271.
- Rasmussen, R.A., 1972. What do the hydrocarbons from trees contribute to air pollution? *J. Air Pollut. Control Assoc.* 22 (7), 537–543.
- Rindelaub, J.D., Wiley, J.S., Cooper, B.R., Shepson, P.B., 2016. Chemical characterization of  $\alpha$ -pinene secondary organic aerosol constituents using gas chromatography, liquid chromatography, and paper spray-based mass spectrometry techniques. *Rapid Commun. Mass Spectrom.* 30 (13), 1627–1638.
- Ripperton, L.A., Jeffries, H.E., White, O., 1972. Formation of aerosols by reaction of ozone with selected hydrocarbons. *Adv. Chem.* 113, 219–231.
- Ryzhkov, A.B., Ariya, P.A., 2004. A theoretical study of the reactions of parent and substituted Criegee intermediates with water and the water dimer. *Phys. Chem. Chem. Phys.* 6 (21), 5042–5050.
- Singh, H.B., Ludwig, F.L., Johnson, W.B., 1978. Tropospheric ozone: concentrations and variabilities in clean remote atmospheres. *Atmos. Environ.* 12 (11), 2185–2196.
- Slade, J.H., Knopf, D.A., 2014. Multiphase OH oxidation kinetics of organic aerosol: the role of particle phase state and relative humidity. *Geophys. Res. Lett.* 41 (14), 5297–5306.
- Song, C., Zaveri, R.A., Shilling, J.E., Alexander, M.L., Newburn, M., 2011. Effect of hydrophilic organic seed aerosols on secondary organic aerosol formation from ozonolysis of  $\alpha$ -pinene. *Environ. Sci. Technol.* 45 (17), 7323–7329.
- Stewart, D.J., Almbrook, S.H., Lockhart, J.P., Mohamed, O.M., Nutt, D.R., Pfrang, C., Marston, G., 2013. The kinetics of the gas-phase reactions of selected monoterpenes and cyclo-alkenes with ozone and the NO<sub>3</sub> radical. *Atmos. Environ.* 70 (2), 227–235.
- Tobias, H.J., Ziemann, P.J., 2001. Kinetics of the gas-phase reactions of alcohols, aldehydes, carboxylic acids, and water with the C13 stabilized Criegee intermediate formed from ozonolysis of 1-tetradecene. *J. Phys. Chem. A* 105 (25), 6129–6135.
- Vartiainen, M., Mcdow, S.R., Kamens, R.M., 1994. Water uptake by aerosol particles from automobile exhaust and wood smoke. *Chemosphere* 29 (8), 1661–1669.
- Witkowski, B., Gierczak, T., 2014. Early stage composition of SOA produced by  $\alpha$ -pinene/ozone reaction:  $\alpha$ -Acylxyhydroperoxy aldehydes and acidic dimers. *Atmos. Environ.* 95, 59–70.
- Witter, M., Berndt, T., Boge, O., Stratmann, F., Heintzenberg, J., 2002. Gas-phase ozonolysis: rate coefficients for a series of terpenes and rate coefficients and OH yields for 2-methyl-2-butene and 2,3-dimethyl-2-butene. In: *J. Chem. Kinet.* 34 (6), 394–403.
- Yasmeen, F., Vermeylen, R., Szmigielski, R., Iinuma, Y., Böge, O., Herrmann, H., et al., 2010. Terpenylic acid and related compounds: precursors for dimers in secondary organic aerosol from the ozonolysis of  $\alpha$ - and  $\beta$ -pinene. *Atmos. Chem. Phys.* 10 (19), 9383–9392.
- Zhang, H., Hu, D., Chen, J., Ye, X., Wang, S.X., Hao, J.M., et al., 2011. Particle size distribution and polycyclic aromatic hydrocarbons emissions from agricultural crop residue burning. *Environ. Sci. Technol.* 45 (13), 5477–5482.
- Zhang, X., Mcvay, R.C., Huang, D.D., Dalleska, N.F., Aumont, B., Flagan, R.C., Seinfeld, J.H., 2015. Formation and evolution of molecular products in  $\alpha$ -pinene secondary organic aerosol. *Proc. Natl. Acad. Sci. U. S. A.* 112 (46), 14168–14173.